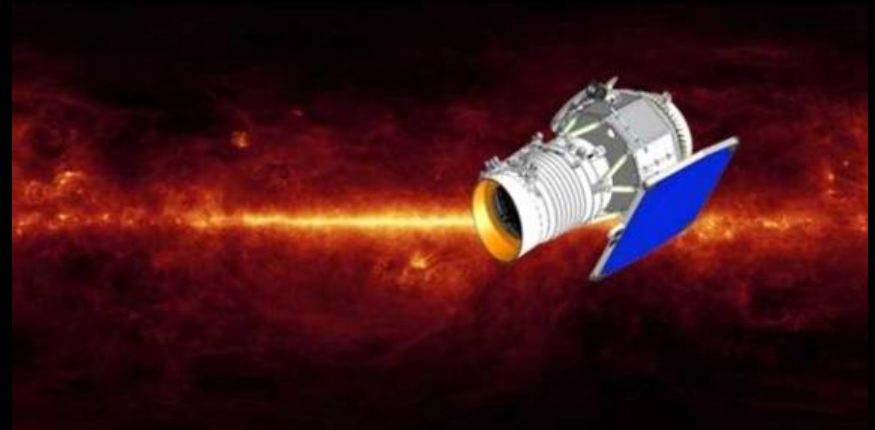
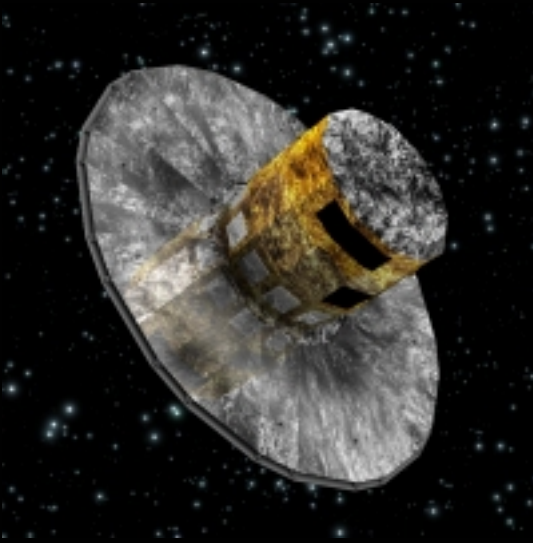


# Asteroid Surface Ages (and Diameters, and Volumes, and Albedos) from Gaia and WISE



**Michael “Migo” Mueller, Marco Delbo’ (OCA Nice, FR),  
Josef Ďurech (Charles University Prague, CZ)**

# Overview

**Gaia: 10,000—100,000's shapes and spin axes of asteroids**

- **Allows better understanding of thermal data**
- **Better constraints on  $D$  and  $p_V$** 
  - Itokawa, Lutetia, and friends
  - Mass densities to within 30—40% or so
- **Thermal inertia**

**What is thermal inertia and why should I care?**

- **Yarkovsky/YORP, regolith, surface age**

**Data situation**

- **IRAS, Spitzer, WISE, ...**

**Modeling situation**

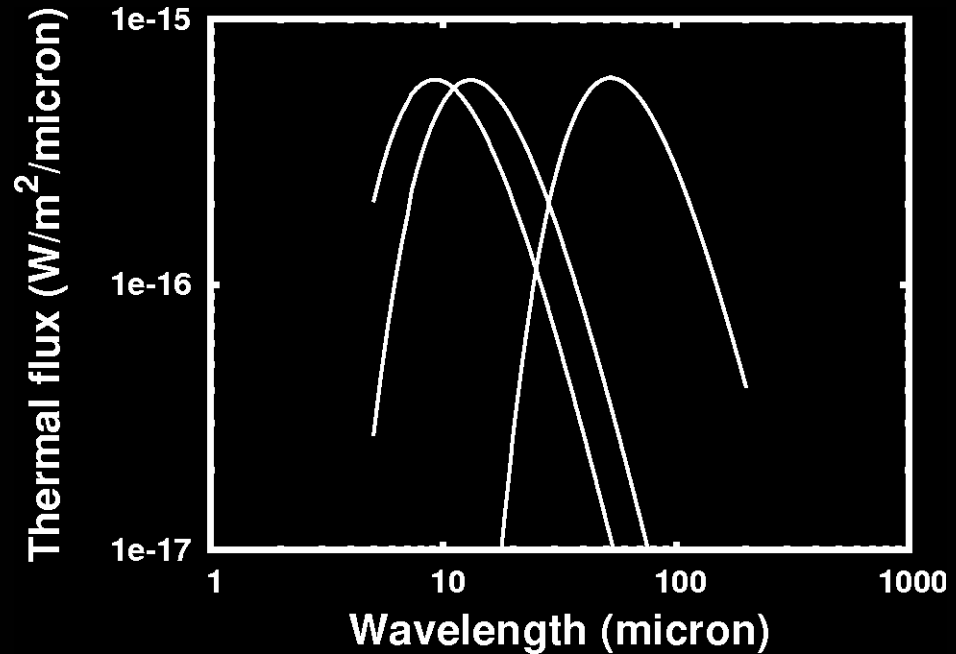
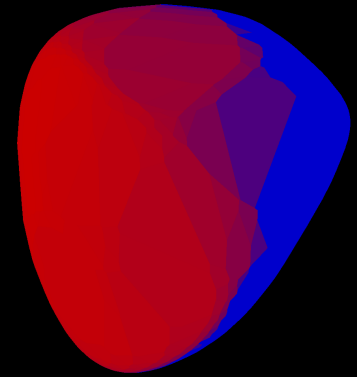
**Precursor study: IRAS + DAMIT**

**Outlook: WISE + Gaia**

# Thermal Emission of Asteroids

$$F(\lambda) = D^2 \int \text{Planck}(\lambda, T(A)) dA$$

- Peaks between  $< 10 \mu\text{m}$  (NEAs) and  $\sim 100 \mu\text{m}$  (TNOs)
  - Observationally difficult but doable
- Need adequate thermal model
- Good handle on size and albedo
  - (the latter with optical photometry = measure of  $D^2 \cdot p_v$ )
- ... and on any thermal property

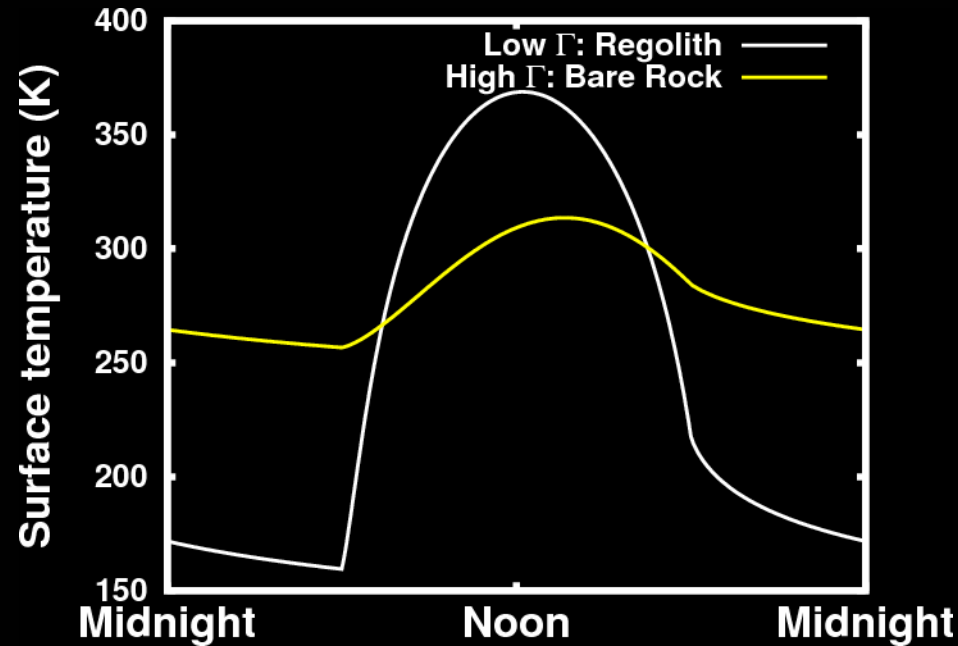



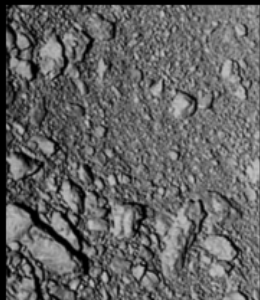
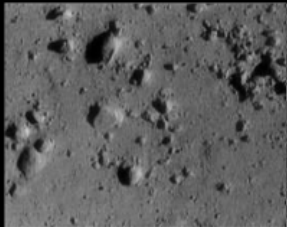

# Thermal Inertia $\Gamma$

Resistance to change in surface temperature

- $\Gamma = \sqrt{\rho c k}$
- Triggers Yarkovsky effect
- Important for mission planning

Regolith properties / surface age from  $\Gamma$



Bare rock $\Gamma \sim 2500$	Itokawa $\Gamma \sim 700$	Eros $\Gamma \sim 150$	The moon $\Gamma \sim 50$
			
	Coarse regolith; pebbles	Immature regolith	Mature, fine regolith

# What Do We Currently Know About $\Gamma$ ?

Regolith even on small NEAs.

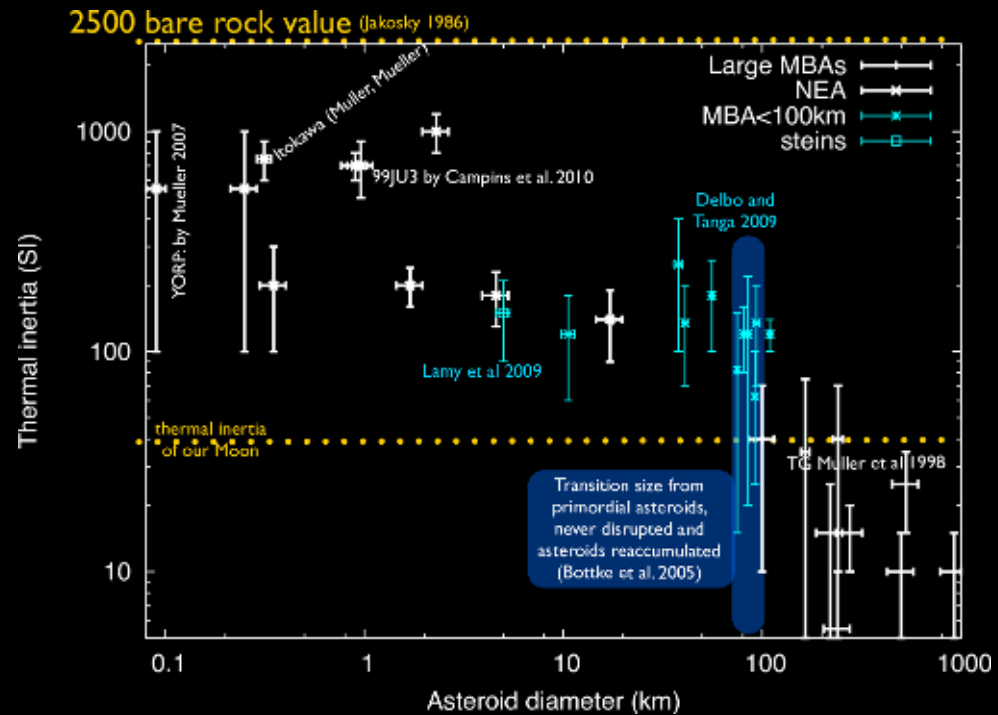
- How does it form?  
How does it stick?

Hints at correlation with D  
(Delbo' et al., 2007)

- Important for Yarkovsky and YORP
- Mission planning

Any dependence on taxonomy?  
On orbit?  
Family?  
Anything?

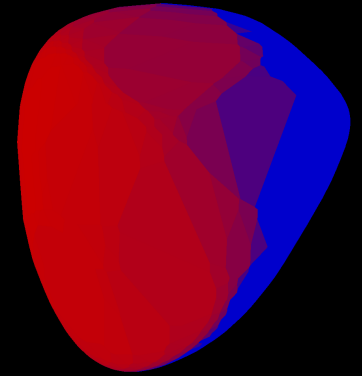
We need more data!



# Thermophysical Model (TPM)

Takes direct account of

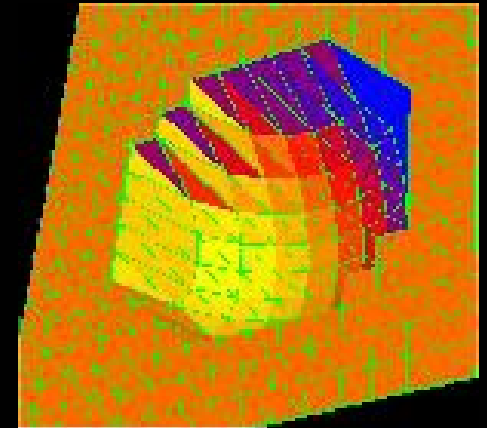
- Shape and spin state
- Thermal conduction leading to thermal inertia
- Surface roughness



References: e.g., Spencer (1990), Lagerros (1998), Mueller (2007)

$$\frac{\partial}{\partial t'} T = \frac{\partial^2}{\partial z'^2} T$$

$$J_{\text{abs}} = \epsilon \sigma T^4 - \Gamma \sqrt{\omega} \frac{\partial}{\partial z'} T$$



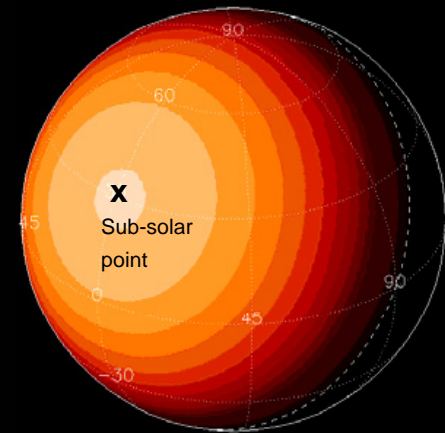
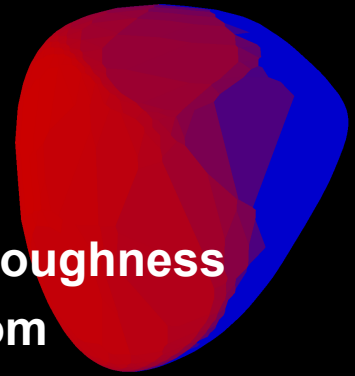
# Two Modeling Approaches

## Thermophysical modeling (TPM)

- Detailed model of shape, spin, thermal conduction, roughness
- Need: high-quality multi-wavelength thermal data from multiple nights / epochs; **model of shape + spin state**
- **Thermal inertia**
- For now: only doable for ~single objects at a time

## Simple models (STM, NEATM)

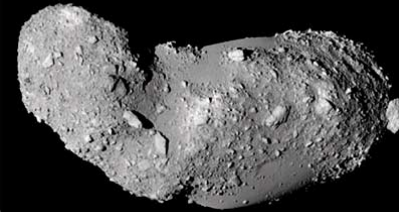
- Simplistic assumptions on shape, temperature
- Less unknowns, less data required
- $\Delta D \sim 15\text{—}20\%$ ,  $\Delta p_v \sim 30\text{—}40\%$   
from modest-quality single-night data
- Most prolific source of asteroid D and  $p_v$
- Great for statistical studies / coarse studies of single objects



# TPM Diameters: “Ground Truth”

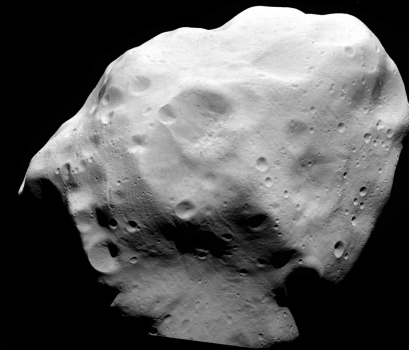
(25143) Itokawa: Hayabusa rendezvous target

- D from Th. Müller et al. (2005):  $0.32 \pm 0.03$  km
- D from Demura et al. (2006):  $0.327 \pm 0.006$  km



(21) Lutetia: Rosetta flyby target

- D from M. Mueller et al. (2006):  $98.3 \pm 5.9$  km
- D from Rosetta flyby:  $\sim 98$  km



We estimate TPM diameters to be within 10%,  
*that's probably conservative* (more work needed).

**Mass densities from AO + TPM can be good within 30—40 %** – if shape + spin axis are known, if good thermal data are available, and if thermal modeling is done properly (e.g., Mueller et al., 2010).



# Thermal Data Situation

## Ground-based observations:

- Extremely challenging due to high background
- Only practical for select targets

## Current largest data set from IRAS

- All-sky survey in 1983 (!) at 12, 25, 60, 100  $\mu\text{m}$
- Coarse  $D$  and  $p_v$  of 2,288 (bright) asteroids
- Latest analysis: *Ryan & Woodward (2010)*



## WISE data have just been taken (2010—2011)

- All-sky survey at 3.6, 4.5, 12, and 25  $\mu\text{m}$
- 100,000's of asteroids
- Waiting for first useful calibrated data release...
- Nominal data analysis based on assumed spherical shape
- **With Gaia shapes, we can do better!**



# Precursor Study: IRAS + DAMIT

Can we really run the TPM in batch mode? Fully automatically, w/o crashes? Within reasonable CPU time? Do results make any sense?

Precursor study: IRAS fluxes + DAMIT shapes and spin state

- Download IR fluxes, shapes, and spin states automatically
  - 94 objects as of 28 Oct. 2010
- Retrieve ephemeris, correct for lighttime + filter breadth
- Run TPM
- Gather results, estimate error bars

Result:

- IDL batches run smoothly w/o crash nor human interference
- TPM runs done within ~30 days
- $D$ ,  $p_v$ , and  $\Gamma$  results of previous studies are reproduced (Mueller et al., 2006; Delbo' & Tanga, 2009 for the most part...)

# What Does This Mean for Gaia + WISE Data Set?

TPM analysis can be done in ~300 days

- Using my current computer (...) but after tweaking TPM code for speed (factor 10+ speed-up possible)

Will provide best catalog of asteroid  $D$ ,  $p_v$ , and  $\Gamma$  (10,000+ objects)

- Synergy between Gaia and WISE, and maybe ExploreNEOs

Increase # of measured  $\Gamma$  by 3 orders of magnitude!!!

- Allows first systematic studies of
  - Regolith vs taxonomy
  - Regolith vs orbit (inner vs outer belt, ...)

Long-term goal: correlate  $\Gamma$  with surface age, constrain dynamical + collisional history of asteroid main belt

- Observational handle:  $\Gamma$  of collisional families of known age
- Auxiliary work: lab studies of thermal inertia of regolith particles; theoretical modeling of regolith formation / aging